



The Llamas de Cabrera gold district, a new discovery in the Variscan basement of northwest Spain: A fluid inclusion and stable isotope study

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ABSTRACT

The northwest of the Iberian Massif is well known to host numerous gold mineralizations which were exploited during Roman times. This study presents a multidisciplinary approach leading to the mineralogical, fluid inclusions, and geochemical characterization of the recently rediscovered Roman gold district of Llamas de Cabrera. Gold occurs in extensional quartz veins hosted in the upper layers of the Lower to Middle Ordovician Armorican quartzite and occasionally in the overlying Middle Ordovician slates from Luarca Formation. The veins are characterized by multistage ore deposition developed along three hydrothermal stages: As–Fe (I), As–Fe (II), and Au–Zn–Cu–Pb, followed by late supergene alteration processes. The fluid inclusion study undertaken on mineralized quartz revealed the presence of three types of fluids: (1) CO₂–(CH₄) fluid inclusions, (2) aqueous-carbonic fluid inclusions dominated by CO₂ and subordinate CH₄ related to quartz–arsenopyrite–pyrite deposition, and (3) aqueous fluid inclusions related to the Au-sulfide deposition. In conjunction with arsenopyrite geothermometry and fluid-inclusion data an attempt was made to determine the P–T conditions of ore formation. Changes occurred in the P–T conditions from arsenopyrite and pyrite deposition in quartz veins from aqueous-carbonic fluids at 300–390 °C and 200–220 MPa towards 180–310 °C and <200 MPa at the stage of gold from aqueous fluids. Mixing of two aqueous fluids of relatively contrasted salinity were favorable factors for decreasing gold solubility and could allow gold deposition. The δ³⁴S values are similar for the two As–Fe stages, +8.0‰ to +16.3‰ and +9.0‰ to +19.5‰ respectively, and for the post-S₁ pyrites from siliciclastic Luarca Formation suggesting a comparable sulfur source. Combining calculated δ¹⁸O values of fluids of +6.7 to +9.4‰ with fluid inclusion data indicates that fluids of unknown origin are largely equilibrated with metamorphic lithologies at medium–high temperatures. The ascent of some late Variscan magmatic bodies, not exposed at the present day erosion level, provides the heat source for convective water circulation. Gold quartz veins from Llamas de Cabrera display similar features and processes to related deposits of similar type elsewhere in the Variscan belt of western Europe (northwest and central/western Iberia, French Massif Central, Bohemian Massif). However Llamas de Cabrera shows certain differences, such as the non-appearance of igneous rocks in the area and the absence of Sb-rich minerals compared to some Variscan Sb–Au deposits, including northern, central, and western Iberia. This last could probably be due to a greater depth of formation for the As–Au veins in comparison to Sb–Au veins.

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1. Introduction

The Iberian Massif is the southwest extension of the European Variscan Belt and one of the largest domains of the Variscan orogen. The northwest of the Iberia Massif was one of the most important

gold producing areas during Roman times. Two types of mineralization, gold-quartz veins and alluvial deposits, were exploited between the first and third centuries when the region was part of the Roman Empire. Studies of these gold deposits have been made focusing on the archaeological, geological, mineralogical, and geochemical aspects of the mineralization (Arias et al., 1997; Boiron et al., 1996; Crespo et al., 2000; Herial, 1984; Noronha et al., 2000; Ribera et al., 1992; Tornos et al., 1997), suggesting gold lodes genesis is related to convective hydrothermal cells re-equilibrated with metamorphic rocks. The district of Llamas de Cabrera has not previously been studied in detail. Although it was

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extensively mined by Roman miners it remained unknown until 2002, when it was rediscovered by Matias and Gómez-Fernández (2003).

Mining in Llamas de Cabrera district started in ancient times. The spectacular Roman mine workings in the northwest of the Iberian Peninsula have been the subject of several archeological studies (Pérez-García et al., 2000; Sánchez-Palencia, 1983). About 500 historical mining sites have been recorded in this area and the total gold production during the first and second centuries has been established to be around 195 t, mainly from alluvial types, with an average gold grade of 67 mg/m³ and maximum of 1–2 g/m³. From Roman times until the present, the gold mines in northwest Spain have remained inactive, except for small operations in the late nineteenth and first half of the twentieth centuries. However, since the increasing of gold prices in the 1970s, exhaustive prospecting was developed. The Llamas de Cabrera district shows substantial ancient mine working, open pits, vertical pits, and galleries, with its own hydraulic network system (Gómez-Fernández et al., 2005; Matias and Gómez-Fernández, 2003).

The primary gold ores of Llamas de Cabrera are located in locally post-deformation quartz veins hosted in the transition zone between the Armorican quartzite and the overlying black slates (Lower to Middle Ordovician). These deposits could be classified as the so-called “mesothermal” or “mesozonal” sulfide-poor gold-bearing quartz veins associated with regionally metamorphosed terranes of all ages (Groves et al., 1998). Widespread similar examples are found in Europe in relation to the Variscan orogen, for example the Massif Central of France (Boiron et al., 2003; Bouchot et al., 2005; Essaraj et al., 2001) and the Bohemian Massif (Boiron et al., 2001). These deposits were considered as orogenic vein gold deposits hosted in Late Proterozoic–Early Paleozoic metamorphic rocks and emplaced between 340 and 300 Ma. The northwest of the Iberian Peninsula gives a good opportunity to compare the characteristics of Au–As veins with Au–Sb ore deposits located in Northwest and Central Iberia.

The present contribution (1) provides the first systematic description of Llamas de Cabrera gold district, (2) characterizes the hydrothermal ore-forming system through detailed fluid inclusion study, (3) analyses the compositional variation of arsenopyrite in order to be used as a geothermometer, (4) deduces information on possible sources of the mineralizing fluids and materials through stable isotope study, (5) reconstructs the P–T path of the history of the deposit, and (6) improves the knowledge of the auriferous veins of the northwest of the Iberian Peninsula. Finally, this paper synthesizes the main characteristics of the gold-bearing deposits of the Iberian Massif and compares to other similar deposits in western Europe.

2. Geological setting

According the division of the Iberian Variscan Massif (IVM) proposed in Pérez-Estaún et al. (2004), the studied area belongs to the Central Iberian Zone (CIZ), just in the limit to the West Asturian Leonese Zone (WALZ) (see Fig. 1.a inset). The gold mining works are placed along the northern limb of the Truchas Synform (Fig. 1.a) a complex structure formed by the coaxial interference of Variscan phase 1 and phase 3 events and several north-verging phase 2 thrust structures (Marcos, 1973; Pérez-Estaún, 1975; Rodríguez Fernández et al., 1982). Some of the complexities of this north limb are related to tight sub-vertical Variscan folds developed in Ordovician low grade metasediments and the underlying Cambrian–Ordovician igneous complex. These folds display axial plain related tectonic foliation showing that they were developed during Variscan phase 1. Some trivial continental deposits (not represented in Fig. 1.a), related to the current fluvial net, quaternary glacial episodes, and slope avalanches, blanket the region locally. Nevertheless, some of these thin sedimentary bodies host quantities of gold and focused the mining works during Roman times (Herail, 1984; Matias and Gómez-Fernández, 2003; Pérez-García et al., 2000).

The Palaeozoic sedimentary record spans all of the Ordovician System, displaying one of the most complete sequences in the CIZ. Regionally, the Ordovician rocks overlie a voluminous igneous unit which crops extensively out to the South and East of the zone. This acidic igneous body is the oldest rock exposed in the region and is named “Ollo de Sapo” (Toad’s Eye) Formation after a local name. It has been interpreted as a group of magmatic, volcanic, and volcano-sedimentary rocks of uppermost Cambrian and low Ordovician age (Díez Montes et al., 2010). It is gradually overlaid by a thick succession of phyllite and quartzite which becomes more enriched in the sandy term towards the top, ending in a decametre-thick bed of white quartzite (Armorican type). Trace fossils found in the area validate the Lower Ordovician age of this formation (Martínez Catalán et al., 1992). The Middle Ordovician is made of a characteristic unit of black slates. The transition between the Armorican quartzite type levels and the black slates is gradual, and some authors distinguish a transitional series locally, called Rubianes Formation in the region (Gutiérrez-Marco et al., 1999), which is formed by layers of quartzite, slate, and oolitic iron and hosts most of the mineralized bodies studied here. The Upper Ordovician sequence lies unconformably and starts with the “La Aquiana” limestone. The rest of the Upper Ordovician was divided by Barros Lorenzo (1989) into three detrital formations: Casao, Rozadais, and Losadilla from bottom to top.

The poly-phase Variscan tectono-metamorphic processes affected the entire IVM. In the Truchas Syncline North limb no other tectonic foliation than the Variscan phase 1 was identified, suggesting that the mineralized veins are younger than this foliation. Variscan magmatic bodies of any phase are not exposed in the work area. But towards the south and east, in the “Ollo de Sapo” antiform, diverse Variscan granitoides are exposed. They include early syn-kinematic, syn-kinematic mesozonal, syn-kinematic epizonal, and post-kinematic epizonal granitoides related to a metamorphic core complex (Díez Montes et al., 2010). At present-day surface level the grade of regional metamorphism is low (chlorite zone) in the Ordovician sequence rocks. According Díez Montes (2007) the metamorphism is coeval to the Variscan phase 1, reaching its peak towards the end of this phase. The same author established a different metamorphic history for the “Ollo de Sapo” formation (structurally deeper during the Variscan) with a higher metamorphic peak during the late Phase 2 and a retrograde subsequent evolution related to the thermal doming processes associated to the formation of a metamorphic core complex.

3. Vein features

Llamas de Cabrera gold district constitutes an extensive area (45 × 20 km) which contains several occurrences of quartz-arsenopyrite gold veins (Fig. 1). The ore deposits are structurally controlled and could be grouped in three types (Table 1): (1) N–S to N 20° E extensional quartz veins, (2) quartz veins, brecciated structures, and stockworks along WNW–ESE, and (3) other ore are related to NE–SW faults. They are hosted in the upper layers of the Lower to Middle Ordovician Armorican quartzite facies (Fig. 3a), equivalent to Rubianes Formation. Occasionally mineralized quartz veins have been likewise recognized in Middle Ordovician slates from the Luarca Formation (Fig. 3b).

The highest density of quartz veins appears in the western part of the district, in the Llamas de Cabrera area, within a 75 × 450 m E–W segment. The set of sub-parallel veins displays preferably N–S orientation (Zone B and C), as well as N 124° E in Zone A with a sub-vertical dip ranging from 75 to 85°. The veins filled by milky quartz are up to 0.5 m thick. These veins show all the characteristics of dilational veins, and especially do not show any shearing of fracture fill. Mineralized tectonic breccias structures are characteristics of major WNW–ESE and NE–SW faults in A Zone and in the Pombriego deposit (Fig. 3c). The Valdecorrales area displays mining activity along 325 m trending N 70° E. Pombriego and Tabuyo del Monte show minor mining activity constituting narrow mineralized areas along N 120° E and N 130° E

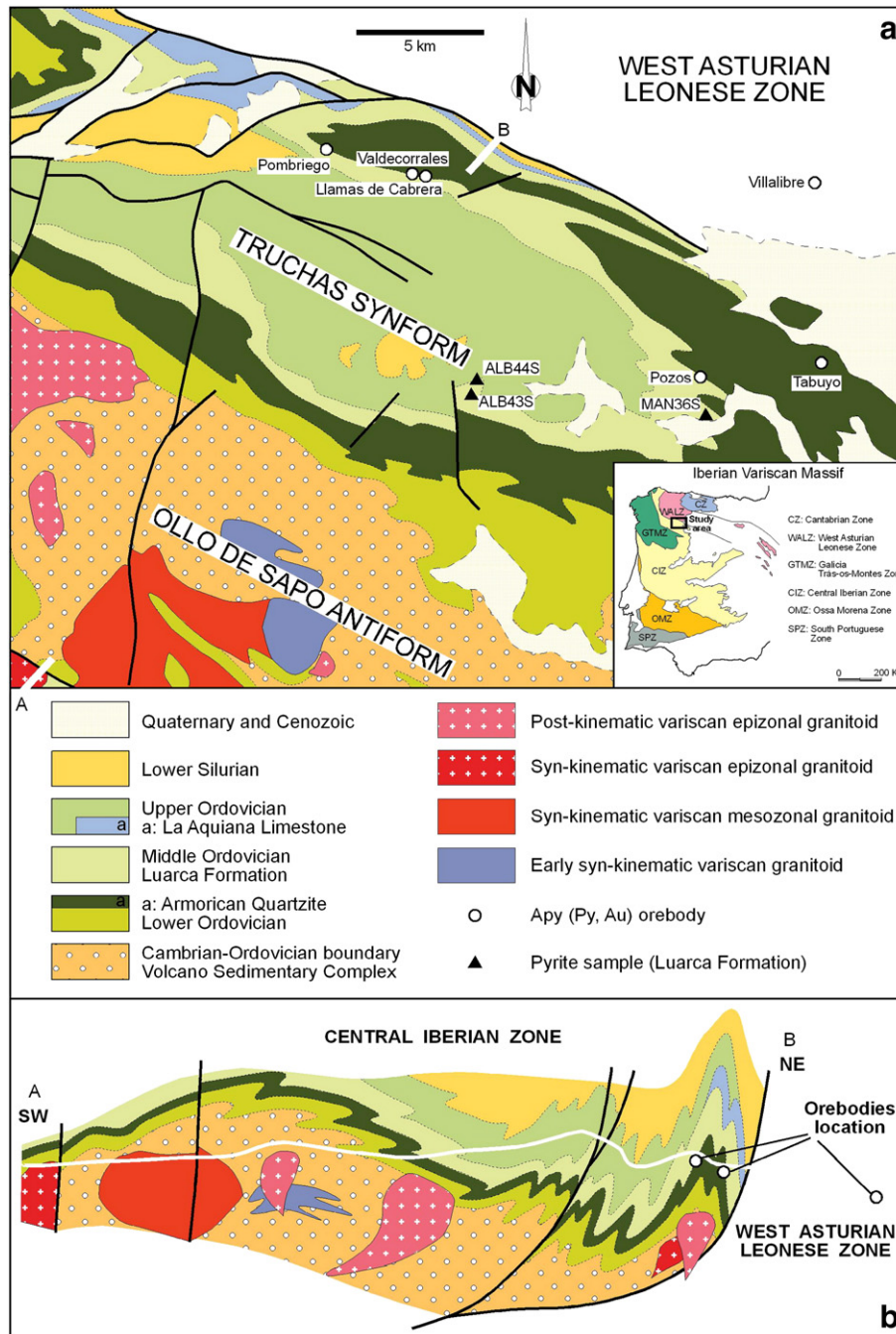


Fig. 1. a: Regional geological sketch displaying part of the Ollo de Sapo Antiform and the Truchas Synform, in the NE of the Central Iberian Zone of the Iberian Variscan Massif (see the inset for general location). The studied ore bodies are shown as open circles. Location of pyrite samples from the Middle Ordovician unit are shown as triangles. b: Interpretative cross section along the AB line which both ends are displayed in white on A draw. The presence of intrusive bodies in the SW was inferred from the outcrops. NE granitic intrusions were deduced as the plausible thermal source necessary for the mineralization systems.

faults respectively. Pozos is characterized by the presence of unexploited stockworks on both sides of a major N 120° E fault zone (Table 1).

Hydrothermal alteration, including silicification, chloritization, and sericitization, have been observed in the enclosing rocks. Disseminated mineralization is present in hydrothermally altered zones on both sides of the veins (± 1 m).

4. Sampling and analytical procedures

For this study, samples of quartz and ore minerals from several deposits of Llamas de Cabrera district: Valdecorrales, Llamas de Cabrera A, B, and C, Pombriego, Villalibre, Pozos, and Tabuyo del Monte, were

collected on surface mine exposures. The characteristics of these veins are summarized in Table 1.

Compositional analyses of sulfides (arsenopyrite, pyrite, bismuthinite, sphalerite, chalcopryite, pyrrothite, tetrahedrite, galena) were carried out using an electron microprobe (JEOL Superprobe JXA-8900 M) employing wavelength dispersive spectrometry at the Luis Brú Microscopy Centre (Universidad Complutense de Madrid). The diameter of the electron beam was 5 μ m and all determinations were performed at a potential of 50 kV for sulfides and a beam current of 30 nA. The following minerals were used as standards: galena, anglesite, nickeline, AsGa, HR-160 (Co–Ni–Cr), chalcopryite, Zn, Mo, Ag, Cd, Sb, Bi₂Te₃, cinnabar, and hutchinsonite.

Table 1
Main geological and mineralogical features of gold veins, Llamas de Cabrera district (from W to E).

Deposit	Vein location	Orientations of veins	Country rock	Ore mineralogy	Other characteristics
Pombriego		Along a N120°E/74°S fault	Quartzite (Lower to Middle Ordovician)	Qz, Apy, Py, Po, Gn, Ccp, Au, Ttr	Tectonic breccia
Llamas de Cabrera	Valdecorrales	Along a N70°E fault	Quartzite (Lower to Middle Ordovician)	Qz, Apy, Py, Gn, (Au)	Veinlets and tectonic breccias related to a NE–SW fault.
	A Zone	N124°E/75°N	Quartzite, sandstone and slate (Lower to Middle Ordovician)	Qz, Apy, Py, Sp, Gn, Au, Ccp, Po, Bis	Related to NW–SE “La Patadura” fault. Tectonic breccia
	B Zone (El Zanjón)	N175°E/85°N			Individual quartz vein
	C Zone	N–S/80°E			More than twenty N–S mineralized quartz veinlets E–W aligned
Pozos		Along a N120°E/80–90° fault	Quartzite and slate (Lower to Middle Ordovician)	Qz, Apy, Py, Au, Sp, Gn, Ccp, Po, Ttr, Cin*	Stockwork
Ermita de Villalibre		N18°E/80°W	Quartzite and slate (Middle Cambrian to Lower Ordovician)	Qz, Apy, Py, (Au)	Quartz veinlets aligned WNW–ESE
Tabuyo del Monte		N145°E/40°N and N130°E/50°S	Quartzite (Lower to Middle Ordovician)	Qz, Apy, Py, (Au)	Set of quartz veinlets

Ore Mineralogy: primary minerals. Qz: quartz, Apy: arsenopyrite, Py: pyrite, Sp: sphalerite, Gn: galena, Ccp: chalcopyrite, Po: pyrrhotite, Au: gold, (Au): inferred, not visible, Bis: bismuthinite, Ttr: tetrahedrite, *cin: cinnabar from Tornos et al. (1997).

Microthermometric studies of fluids inclusions (86 from 15 samples) were carried out on doubly polished wafers (300 μm) with a Linkam THMSG 600 heating–freezing stage (Shepherd, 1981). The stage was calibrated based on the melting point of solid standards at $T > 25^\circ\text{C}$ and natural and synthetic inclusions at $T < 0^\circ\text{C}$. The rate of heating was monitored in order to obtain measurement precision of $\pm 0.2^\circ\text{C}$ during freezing and $\pm 1^\circ\text{C}$ when heating within the 25–400 $^\circ\text{C}$ temperature range. The salinity of H_2O –NaCl inclusions reported as equivalent weight percent NaCl (wt.% NaCl equiv.) was calculated from microthermometric data (ice melting, $T_{\text{m,ice}}$) using the equations from Bodnar (1993). Bulk composition and molar volumes were computed from P–V–T–X properties of individual inclusions in the C–O–H–(N–S) system (Bakker et al., 1996; Thiery et al., 1994) using an estimate of the NaCl content but neglecting the effect of sulfate content. The P–T properties of aqueous–carbonic inclusions were modeled using the equation of state of Bakker (1999) for the H_2O –NaCl– CO_2 – CH_4 system and data from Zhang and Frantz (1987) for the H_2O –NaCl system. The analysis of the gas phase composition of individual inclusions was performed on a Jobin Yvon T64000 multi-channel modular micro-Raman spectrometer following the procedure given by Burke (2001) and analyses of the relative amounts of CO_2 and CH_4 . Analyses were performed at the Serveis Científicotècnics of the Universitat de Barcelona (Spain).

Twenty-seven quartz samples were analyzed for $^{18}\text{O}/^{16}\text{O}$ at the Servicio General de Isótopos Estables of Salamanca University (Spain) using a laser fluorination procedure, involving total sample reaction with excess ClF_3 (Borthwick and Harmon, 1982) using a CO_2 laser as a heat source (in excess of 1500 $^\circ\text{C}$; following Sharp, 1990). This O_2 was then converted to CO_2 by reaction with hot graphite (Clayton and Mayeda, 1963) and then analyzed on-line by a VG-Isotech SIRA-II mass spectrometer. Reproducibility is better than $\pm 0.3\%$ (1σ), based on repeat analyses of internal and international standards during sample runs. Results are reported in standard notation ($\delta^{18}\text{O}$) as per mil (‰) deviations from the VSMOW standard.

Sulfur isotope analyses were performed on 13 arsenopyrites, 8 pyrites, and 3 sphalerites. Analyses were performed at the Servicio General de Isótopos Estables of Salamanca University (Spain) by standard techniques (Robinson and Kusakabe, 1975) in which SO_2 gas was liberated by combusting with excess Cu_2O at 1075 $^\circ\text{C}$, in vacuo. Liberated gases were analyzed on a VG Isotech SIRA II mass spectrometer, and standard corrections applied to raw $\delta^{66}\text{SO}_2$ values to produce true $\delta^{34}\text{S}$ values. The standards employed were the international standards NBS-123 or NBS-127. These gave $\delta^{34}\text{S}$ values of +17.1‰, –31.5‰, and –4.6‰ respectively, with 1σ reproducibility better than $\pm 0.2\%$. Data are reported in $\delta^{34}\text{S}$ notation as per mil (‰) variations from the Vienna Canyon Diablo Troilite (V–CDT) standard.

5. Mineralogy and paragenesis

The mineralogy of the different studied veins is summarized in Table 1 and the paragenetic sequence is given in Fig. 2. The veins are characterized by multistage ore deposition belonging to three primary mineralizing stages and a supergene alteration stage.

5.1. Ore stage I: As–Fe (I)

During the first ore stage fractures affecting the silicified (microcrystalline quartz), sericitized (Fig. 3d), and local chloritized host rocks were filled by quartz, arsenopyrite, and pyrite. Milky quartz (Qz I) forms aggregates of medium-grained crystals with serrate contacts (Fig. 3d). The quartz crystals exhibit undulose extinction of large rounded to irregularly shaped crystals cemented by a mass of fine quartz grains. Quartz I, the main mineral of the veins, contains abundant small (generally $< 5\ \mu\text{m}$) fluid inclusions and hosts all later mineralization. Arsenopyrite I and pyrite I are the earliest sulfides to be formed. Arsenopyrite is the most common sulfide in the veins and it occurs as massive or as sub-euhedral crystal aggregates during stage I. Pyrite I is rare and occurs intergrown with arsenopyrite I.

5.2. Ore stage II: As–Fe (II)

The intermediate stage is characterized by the presence of a second quartz, arsenopyrite, and pyrite generation. A brittle deformation phase is recognizable from cataclastic textures of arsenopyrite (Figs. 3e and 4c), pyrite I, and quartz I, and development of microfractures healed by subsequent ore minerals.

Clear quartz (Qz II) can be observed in the central part of the veins, sometimes having euhedral terminations, and cementing arsenopyrite I (Fig. 3e). The second generation of arsenopyrite II is characterized by euhedral or sub-euhedral rhombs intergrown with pyrite II. No compositional differences were recognized between pyrite I and II (Table 2). Arsenopyrite II is strong zoned with significant variations in As content. Although this compositional zoning cannot be recognized optically, it is easily recognized in back-scattered electron images. Arsenopyrite II is intergrown intimately with a second generation of pyrite (Fig. 4a). Occasionally, scattered pyrite and arsenopyrite porphyroblasts of the second generation can be recognized in slates crosscutting S_1 schistosity. They developed pressure shadows showing syntectonic structures of feather quartz and fibrous muscovite (Fig. 3f). These textural relations indicate the simultaneous equilibrium crystallization of arsenopyrite II with pyrite II. Scarce grains of bismuthinite are recognized in pyrite from this stage.

Stage	Primary mineralization			Supergenic alteration
	I	II	III	
Mineralization	As-Fe (I)	As-Fe (II)	Au-Zn-Cu-Pb	
Quartz	Qz I	Qz II		
Arsenopyrite	Apy I	Apy II		
Pyrite	Py I	Py II		
Bismuthinite				
Native Gold				
Sphalerite				
Galena				
Chalcopyrite				
Pyrrhotite				
Tetrahedrite				
Covellite				
Anglesite				
Scorodite				
Goethite				
Fluid inclusions				
Type 1	----- ?			
Type 2	-----	-----		
Type 3			-----	

Fig. 2. Simplified paragenetic sequence for the assemblages of principal ore and alteration mineral and distribution of fluid inclusion types at Llamas de Cabrera district. Line thicknesses represent schematically the relative mineral abundances. Type 1: CO₂–(CH₄) fluid inclusions, Type 2: aqueous-carbonic fluid inclusions, Type 3: aqueous fluid inclusions.

5.3. Ore stage III: Au–Zn–Cu–Pb

The third stage is characterized by the cogenetic crystallization of gold, sphalerite, galena, chalcopyrite, tetrahedrite, and pyrrhotite filling microfissures crosscutting the earlier stages of quartz and earlier sulfides. Galena is mainly later than sphalerite, as it partially surrounds sphalerite (Fig. 4b). Chalcopyrite contains inclusions of pyrrhotite and tetrahedrite. Microprobe analyses revealed that tetrahedrite could be considered as freibergite because of its relatively high Ag content (24.33 wt.%). Average composition of sulfides minerals from stage III are shown in Table 2.

Native gold is found as anhedral grains, ranging from 1.3 μm to a few micrometers to (long axis) in size, in cracks affecting previous As–Fe sulfides from ore stage I and II (Fig. 4c and d). Coexistence of native gold with other sulfides can be observed (Fig. 4c). Inclusions of gold are rarely found in quartz. Gold composition was determined in a previous study (Gómez-Fernández et al., 2005), with Au ranging between 91.44 and 95.14 wt.%, Ag between 1.99 and 7.62 wt.%, and minor amounts of As (less than 0.08 wt.%), Fe (0.50–0.9 wt.%), S (0.06–0.12 wt.%), Te (0.05–0.14 wt.%), Bi (0.12–0.27 wt.%), and Zn (less than 0.04 wt.%). Cu content is below the detection limit. Quartz associated with gold is insignificant, implying a reduced supply of silica to the system.

5.4. Supergene mineralization

Early mineral assemblages are locally replaced by supergene covellite (Table 2), anglesite, goethite, and especially scorodite replacing arsenopyrite (Fig. 4c and d). This low-temperature alteration assemblage is found in complex sets of micro-cracks that crosscut the primary minerals.

6. Composition of arsenopyrite

Over 80 spot analyses were made on arsenopyrites I and II and several grains were analyzed in the same polished section. Average

composition of arsenopyrites are shown in Table 3. Arsenopyrite grains from Llamas de Cabrera are subtly zoned, showing that the centers of arsenopyrite crystals are S-rich relative to the rims. The existence of zoning reflects minor changing conditions during arsenopyrite growth and attests to the strong resistance of arsenopyrite to internal homogenization. Because of this refractory nature, arsenopyrite composition reflects formation temperature (Kretschmar and Scott, 1976). Arsenopyrite geothermometer should be used with caution (Kerestedjian, 1997; Sharp et al., 1985); however, in conjunction with fluid inclusion data it could provide an estimate of formation temperatures.

The As content of the first generation of arsenopyrite (Apy I) shows a compositional range between 28.3 and 30.9 wt.% and the range of As content of the second generation of arsenopyrite (Apy II) is slightly lower, between 27.2 and 30 wt.% (Fig. 5). Representative values of Fe content are 36.07 and 36.23 wt.% for Apy I and between 35.62 and 36.47 wt.% for Apy II. Total concentration of Ni and Co is small and does not exceed 0.1 wt.% for Apy I and 0.5 wt.% for Apy II. Other elements show similar values for both types of arsenopyrites (Table 3). The slight differences between arsenopyrites from stages I and II imply similar conditions during the formation of the two stages. Arsenic and sulfur are strongly negatively correlated (Fig. 5) following the classical substitution: $\text{FeAs}_{1 \pm x}\text{S}_{1 \pm x}$.

7. Fluid characterization

7.1. Petrography, distribution, and types

Gold quartz veins of Llamas de Cabrera district are devoid of good material for fluid inclusion studies due to the presence of quartz containing only very tiny fluid inclusions. For this reason, systematic sampling of distinct paragenetic stage for detailed fluid inclusion study was not feasible and only some coarsely crystalline quartz I and II were suitable for microthermometric analysis. The distribution of the fluid inclusions is similar in the two types of quartz.

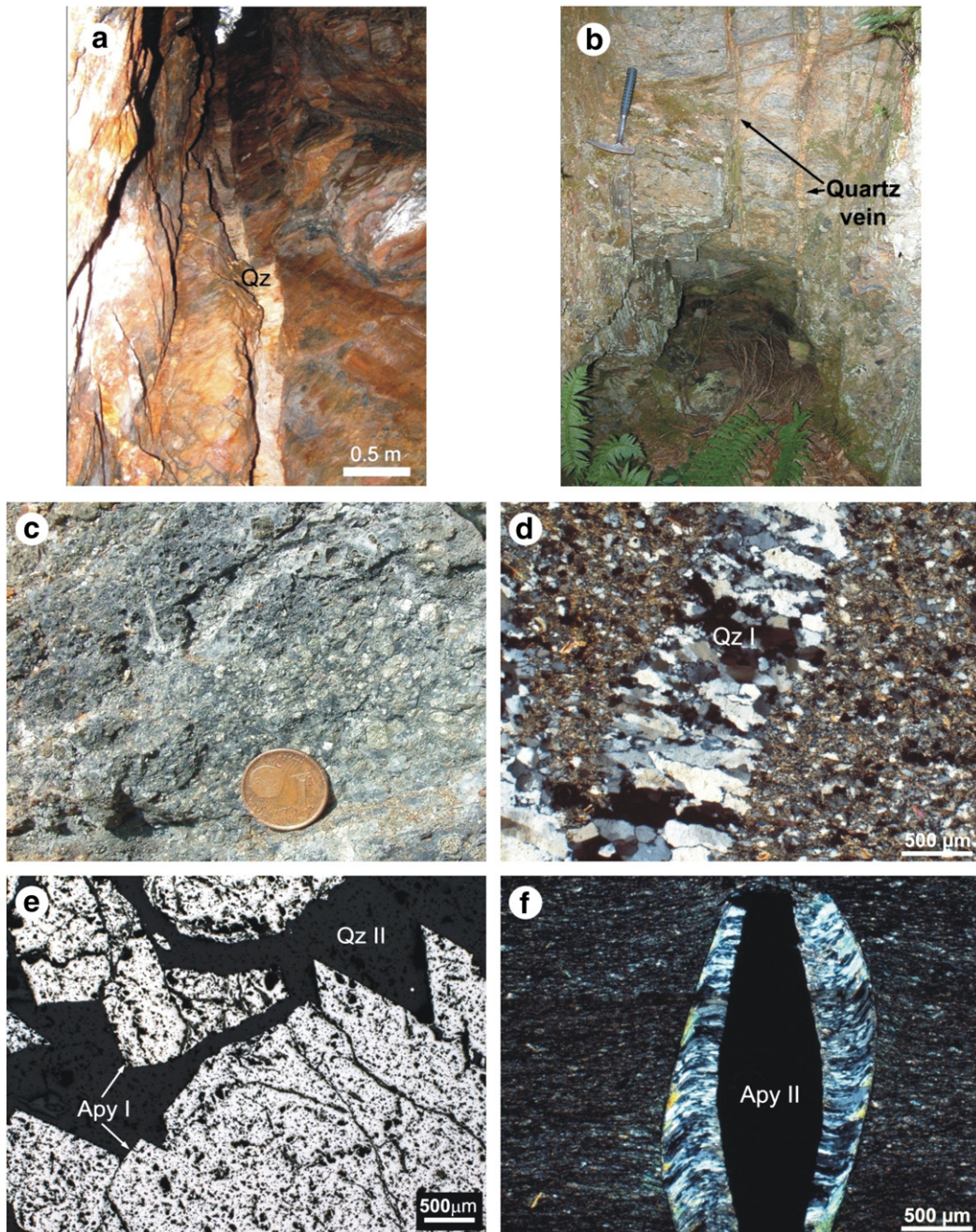


Fig. 3. a) Quartz vein in Armoricain quartzite. Llamas de Cabrera, C Zone b) N–S Quartz vein set in slates from Fm. Luarca. Llamas de Cabrera, C Zone. c) Tectonic breccia. Pombriego. d) Quartz I veinlet in sericitized quartzite, C.N. e) Quartz II filling fractures in arsenopyrite I, 1 N. f) Pressure shadow with quartz and muscovite around a single crystal of arsenopyrite II, C.N. Apy: arsenopyrite, Py: pyrite, Qz: quartz.

Three different compositional types of fluid inclusions are distinguished based on their optical characteristics at room temperature and on microthermometric behavior during cooling and heating experiments. Type 1 (carbonic fluid inclusions) consists of two-phase inclusions lacking visible H₂O at room or sub-zero temperatures, showing a VCO₂ of 0.10 to 0.40 (visual estimate at 25 °C). These inclusions are scarce and have only been recognized in the first quartz generation (Qz I) from Valdecorrales deposit, ranging from 5 to 10 µm in size. Type 2 inclusions H₂O–NaCl–CO₂–(CH₄) are three-phase at room temperature CO₂-bearing inclusions (Fig. 4e), with a dark two fold gaseous bubble and a clear liquid phase with a variable gas/liquid ratio of 0.08–0.40, varying in size between 5 and 15 µm. Type 3 are two-phase (Fig. 4f) and H₂O-rich inclusions (H₂O–NaCl inclusions)

with small vapor bubbles (<30%). These inclusions are abundant and generally relatively small (<5 to 10 µm), showing a variety of shapes ranging from irregular to rounded and ellipsoidal shapes.

Repeated fracturing and healing make it difficult to constrain a precise timing of the different fluid inclusion types relative to the gold event. However, in some cases, detailed petrographic investigations allowed a relative chronology of inclusions to be established using the criteria of Roedder (1984, p.12–45). Most Type 1 and Type 2 inclusions are isolated and regular in shape in quartz (Type I inclusions only in Qz I), suggesting that they could represent early fluids. Presumable re-equilibration of Type I fluid inclusions will be discussed in Section 9. Type 3 are undoubtedly secondary inclusions in relation to the two types of quartz, because they crosscut clusters of

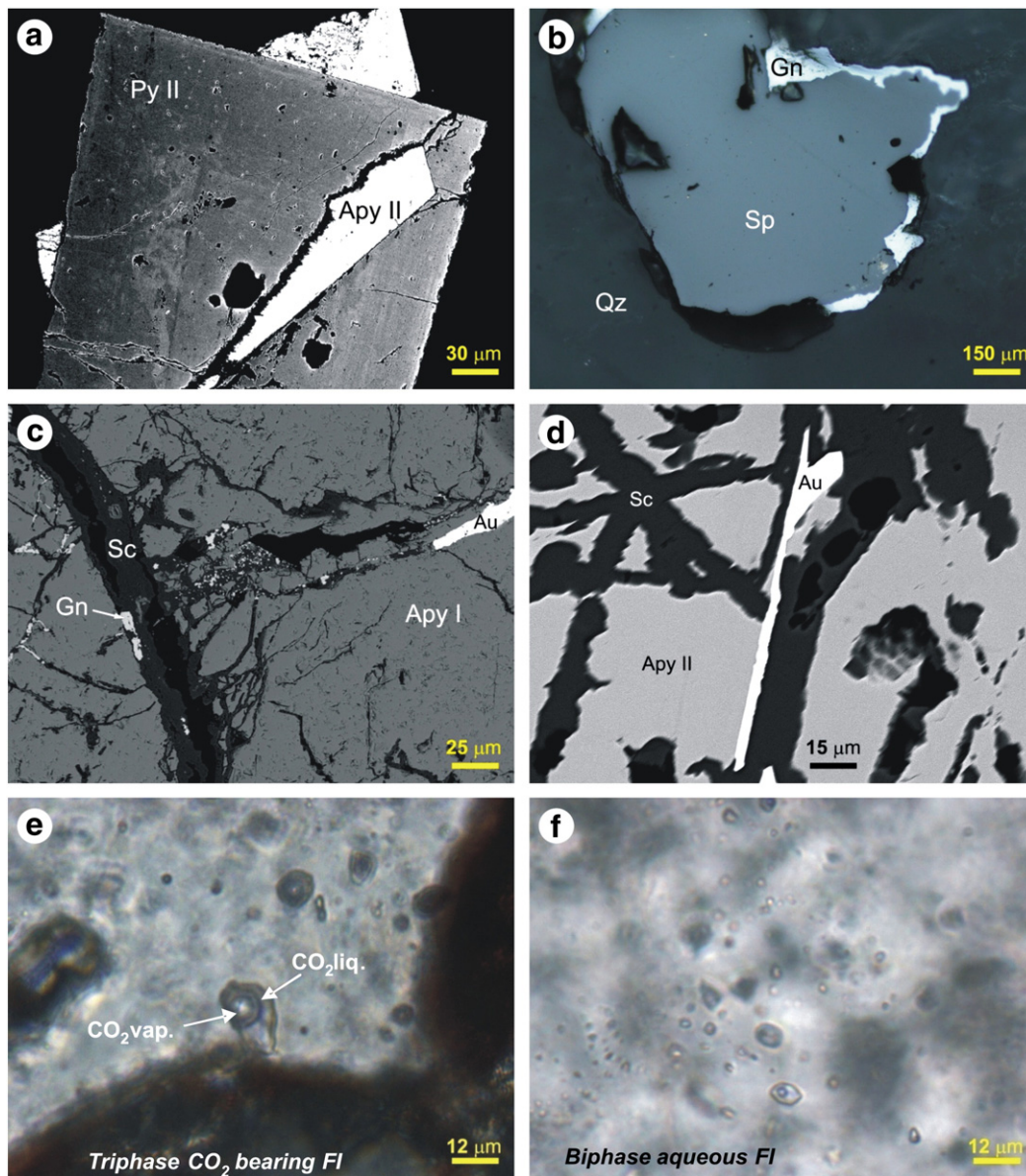


Fig. 4. a) Back-scattered electron (BSE) image of arsenopyrite–pyrite intergrowth. b) Galena surrounding sphalerite, 1 N. c) Back-scattered electron (BSE) image of Au and galena infilling cracks in arsenopyrite. d) Back-scattered electron (BSE) image of Au in fractures in arsenopyrite II. e) Type 2 three-phase aqueous-carbonic fluid inclusions in stage I quartz. f) Type 3 two-phase aqueous fluid inclusions in stage I quartz. Apy: arsenopyrite, Sp: sphalerite, Gn: galena, Sc: scorodite.

earlier aqueous-carbonic fluid inclusions or occur along healed traces of microfractures cutting the quartz boundaries. Type 3 inclusions represent a later aqueous fluid after quartz deposition. Relative chronology between fluid inclusion types and ore deposition is shown in Fig. 2.

7.2. Microthermometric and micro-Raman results

Microthermometric data are summarized in Table 4 with all abbreviations used in the text.

Three different types of fluids have been recognized during the microthermometric work and several runs have been undertaken down to the lowest limit of temperature of the heating–freezing stage ($\sim 180^\circ\text{C}$), in order to detect the presence of CH_4 and/or N_2 . However, all inclusions are frozen between -100 and -120°C , indicating that these compounds are not present in significant amounts.

After heating from these low temperatures, the melting of CO_2 (T_{mCO_2}) in carbonic fluid inclusions (Type 1) occurred in a relatively

narrow interval, between -58.4 and -57°C indicating minor amounts of CH_4 , and the homogenization of carbonic phases (T_{hCO_2}) into liquid occurred from 4.4 to 12.0°C , corresponding to high CO_2 densities of 0.9 to 0.85 g/cm^3 . Formation of clathrates was not observed in type 1 fluid inclusions, which was expected given the absence of visible water in the inclusions.

T_{mCO_2} in aqueous-carbonic fluid inclusions (Type 2) ranges between -58.4 and -56.7°C and T_{hCO_2} between 20.7 and 29.3°C to either liquid or vapor phase (CO_2 densities of 0.62 – 0.76 g/cm^3). The clathrate melting temperature (T_{mCl}) is between 7.8 and 9.3°C , indicating salinities of around 1.4 to $4.1\text{ wt.}\%$ NaCl equiv. Total homogenization of aqueous-carbonic fluid inclusions to liquid occurs between 285 and 360°C (Fig. 6a). The T_{mCO_2} values in carbonic and aqueous-carbonic types, slightly below the triple point of pure CO_2 (-56.6°C), indicate the dominance of CO_2 and the presence of other gases in subordinate amounts. This is confirmed by Raman microprobe analysis, which identified only minor amounts of CH_4 (1 – $4\text{ mol}\%$) in the volatile phase. N_2 and H_2S species have not been

Table 2

Average composition of sulfide minerals from Llamas de Cabrera mine district (Py: pyrite, Gn: galena, Ttr: tetrahedrite, Sp: sphalerite, Po: pyrrotite, Ccp: chalcopyrite, Cv: covellite).

Mineral	Py	Py	Gn	Ttr-Fre	Sp	Po	Ccp	Cv
Stage	I	II	III	III	III	III	III	Superg.
N	15	21	9	4	15	3	5	3
Ag wt.%	<0.01	0.01	0.01	24.33	<0.01	<0.01	0.12	0.91
Au	0.02	0.02	<0.01	<0.01	0.31	0.09	<0.01	<0.01
Bi	0.18	0.16	0.20	0.05	0.05	0.15	0.10	0.12
Cd	0.01	0.01	0.12	1.44	1.29	<0.01	<0.01	<0.01
Cu	<0.02	<0.02	<0.02	20.00	<0.02	0.65	31.85	62.70
S	53.71	52.46	13.58	22.89	33.38	39.89	35.07	34.46
Fe	45.61	45.73	0.01	3.79	7.39	57.91	31.03	<0.01
As	0.03	0.98	<0.02	<0.02	<0.02	<0.02	<0.02	<0.02
Sb	0.05	0.04	0.26	27.20	0.05	0.04	0.04	0.09
Pb	0.27	0.31	85.94	0.45	0.18	0.29	0.21	0.26
Zn	<0.02	<0.02	<0.02	0.18	57.41	1.05	1.13	<0.02
Total	99.88	99.72	100.12	100.33	100.06	100.07	99.56	98.54
Ag at.%	0.00	0.00	0.01	14.47	0.00	0.00	0.05	0.41
Au	0.00	0.00	0.00	0.00	0.08	0.02	0.00	0.00
Bi	0.04	0.03	0.11	0.01	0.01	0.03	0.02	0.03
Cd	0.00	0.00	0.12	0.84	0.56	0.00	0.00	0.00
Cu	0.00	0.00	0.00	20.08	0.00	0.44	23.08	47.58
S	67.12	66.19	50.21	45.63	50.37	53.82	50.36	51.84
Fe	32.72	33.18	0.03	4.31	6.40	44.86	25.58	0.00
As	0.02	0.53	0.00	0.00	0.00	0.00	0.00	0.00
Sb	0.02	0.01	0.26	14.29	0.02	0.01	0.01	0.04
Pb	0.05	0.06	49.17	0.14	0.04	0.06	0.05	0.06
Zn	0.00	0.00	0.00	0.17	42.50	0.69	0.80	0.00

detected. No aqueous-carbonic fluid inclusions have been found at Pozos deposit.

The first melting of ice in Type 3 inclusions was recognized near – 21 °C, indicating the predominance of NaCl among dissolved salts. Type 3 are aqueous low salinity fluid inclusions displaying T_{mice} between – 5.6 and – 1.4 °C that correspond to equivalent salinity of 2.4 to 8.6 wt.% NaCl. Homogenization occurs to the liquid phase (T_h) between 130 and 285 °C, but most of the highest values (230–285 °C) were registered at Pozos deposit (Fig. 6b). Densities

Table 3

Average composition of arsenopyrite analyses from Llamas de Cabrera mine district.

Mineral	Arsenopyrite				
	I	I	II	II	II
Stage	I	I	II	II	II
Site	Tabuyo	Llamas	Llamas	Pozos	Pombriego
N	2	19	32	17	11
Ag wt.%	<0.01	0.01	0.02	0.02	<0.01
Au	0.03	<0.01	0.01	0.01	0.01
Bi	0.04	0.10	0.06	0.06	0.09
Co	0.06	0.03	0.04	0.18	0.32
Fe	36.07	36.29	36.20	35.72	35.80
As	42.56	42.13	40.99	41.02	41.43
S	20.69	21.19	22.01	21.83	21.76
Pb	0.11	0.12	0.12	0.12	0.14
Ni	<0.01	0.01	<0.01	0.34	0.07
Sb	0.07	0.15	0.10	<0.01	0.06
Total	99.63	100.06	99.65	99.45	99.71
Ag at.%	0.00	0.00	0.01	0.01	0.00
Au	0.01	0.00	0.00	0.00	0.00
Bi	0.01	0.03	0.02	0.02	0.02
Co	0.05	0.03	0.03	0.16	0.29
Fe	34.69	34.63	34.40	34.03	34.07
As	30.51	29.97	29.04	29.14	29.39
S	34.66	35.23	36.42	36.24	36.08
Pb	0.03	0.03	0.03	0.03	0.04
Ni	0.00	0.01	0.00	0.31	0.06
Sb	0.03	0.07	0.04	0.00	0.02

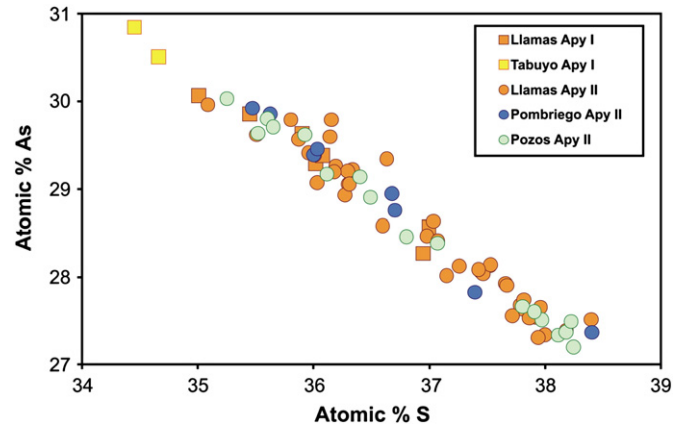


Fig. 5. The content of As and S (at.%) in arsenopyrite (Apy).

range between 0.79 and 0.98 g/cm³. No C–O–(N–S) species have been detected by Raman microprobe in Type 3 fluid inclusions.

8. Stable isotope studies

8.1. Sulfides

Results of sulfur isotope study show that $\delta^{34}S$ values for arsenopyrite, pyrite, and sphalerite are moderately heavy, ranging from + 8.0 to + 23.1‰ with most of the values clustering between + 9.0 and + 14.2‰. The determined range of $\delta^{34}S$ values does not seem to depend on the different analyzed sulfides, arsenopyrite, pyrite, and sphalerite, including arsenopyrite and pyrite from the first and second generation, and does not display significant differences between the deposits (Table 5; Fig. 7). Sulfur isotope analyses were also performed on four post-S₁ pyrites from slates of the Lluarca Formation close to the quartz-gold veins. The $\delta^{34}S$ values of pyrites from slates range between + 7.4 and + 26.3‰, similar to those of sulfides separated from ores.

8.2. Quartz

The $\delta^{18}O$ values for both quartz generations (Qz I and Qz II) from all the studied quartz-gold veins are similar and range from + 9.3 to + 16.1‰. Qz I and Qz II show similar $\delta^{18}O$ values, quartz I between + 9.3 and + 15.9‰ and quartz II between + 13.7 and 16‰. (Table 6; Fig. 8a) Assuming a formation temperature for the two quartz generations (Qz I and Qz II) of between 330 and 290 °C, as deduced from microthermometry of aqueous-carbonic fluids, the calculated $\delta^{18}O_{fluid}$ (using the equation of Zheng, 1993) varied from + 3.2 to + 9.4‰, clustering from + 6.7 to + 9.4‰ (Table 6; Fig. 8b). No suitable fluid inclusions for microthermometric study have been found in Pombriego and no aqueous-carbonic fluid inclusions have been recognized in Valdecorrales and Pozos veins. Therefore $\delta^{18}O_{fluid}$ could not be calculated for Valdecorrales and Pozos samples, VDC-47 and PZ-2, PZ-3 respectively. Estimated temperature using the arsenopyrite composition has been used for $\delta^{18}O_{fluid}$ calculation for quartz from Pombriego.

9. Discussion and conclusion

9.1. Fluid evolution and ore-forming conditions

The appearance of the three types of fluids in Llamas de Cabrera district could be interpreted as an example of continuous fluid evolution, where there was a significant change from CO₂-bearing fluids to dominant aqueous fluids. The ore-forming fluids show an evolutionary trend from early high temperature aqueous-carbonic fluids with low density and salinity to low temperature aqueous fluids with high density and low to moderate salinity. The T_m -salinity plot

Table 4
Microthermometric data from fluid inclusions recognized in quartz. $T_m\text{CO}_2$: melting CO_2 ; $T_h\text{CO}_2$: CO_2 homogenization; $T_m\text{Cl}$: clathrate melting; T_{mice} : last ice melting; T_h : homogenization to vapor (V) or liquid (L); N: number of measurements. N.o.: not observed.

Components/fluid type	Deposit	Sample	N	Vapor (%)	$T_m\text{CO}_2$ (°C)	$T_h\text{CO}_2$ (°C)	$T_m\text{Cl}$ (°C)	T_{mice} (°C)	T_h (°C)	Density (gr/cm ³)	Salinity (wt.% NaCl eq)
CO_2 -(CH_4) Type 1	Valdecorrales	VCD-47 (1)	6	10–40	−58.4/−57.5	4.4/12 L	–	–	–	0.85–0.90	–
	H_2O -NaCl- CO_2 -(CH_4) Type 2	Llamas de Cabrera A	LLA-14	7	20–40	−57.8/−57.1	21.6/22.5 V	8.3/9.1	n.o.	320/332 L	0.74–0.76
Llamas de Cabrera B		LLB-4	6	8–25	−57.6/−57.1	21.8/23.7 L	7.8/8.9	n.o.	306/312 L	0.74–0.76	2.2/4.1
Llamas de Cabrera C		LLC-10–11	6	10–30	−58.4/−57.6	20.7/21.4 L	8.7/9.3	n.o.	307/316 L	0.76–0.77	1.4/2.8
		LLC-8	5	10–30	−57.8/−57.3	22.7/25.2 V	8.8/9.3	n.o.	290/360 L	0.71–0.74	1.4/2.4
Villalibre		VL-1	8	30–40	−57.3/−56.7	25.8/29.3 L	7.9/8.7	n.o.	285/319 L	0.62–0.70	2.6/4.1
Tabuyo		TA-40	5	20–30	−57.3/−57	23.5/27.3 L	8.3/9.0	n.o.	285/295 L	0.67–0.71	2.0/2.4
H_2O -NaCl Type 3	Valdecorrales	VDC-47 (1)	6	10–20	–	–	–	−2.1/−1.4	175/230 L	0.89–0.93	2.4/3.5
	Llamas de Cabrera A	LLA-14	4	10–30	–	–	–	−4/−3.4	185/192 L	0.92–0.93	5.5/6.4
	Llamas de Cabrera B	LLB-4	5	10–20	–	–	–	−2.8/−3.2	187/196 L	0.91–0.92	4.5/4.8
	Llamas de Cabrera C	LLC-1011	5	8–10	–	–	–	−5.6/−2.1	138/211 L	0.92–0.96	3.4/8.6
		LLC-8	4	8–10	–	–	–	−2/−4	135/138 L	0.93–0.96	3.3/6.4
	Pozos	PZ-5(4)	9	8–20	–	–	–	−4.5/−2.9	230/285 L	0.83–0.92	4.7/7.1
	Villalibre	VL-1	4	10–15	–	–	–	−2.3/−1.6	132/140 L	0.84–0.95	2.7/3.9
	Tabuyo	TA-40(1)	6	5–20	–	–	–	−4.7/−1.8	138/276 L	0.82–0.95	2.9/7.4

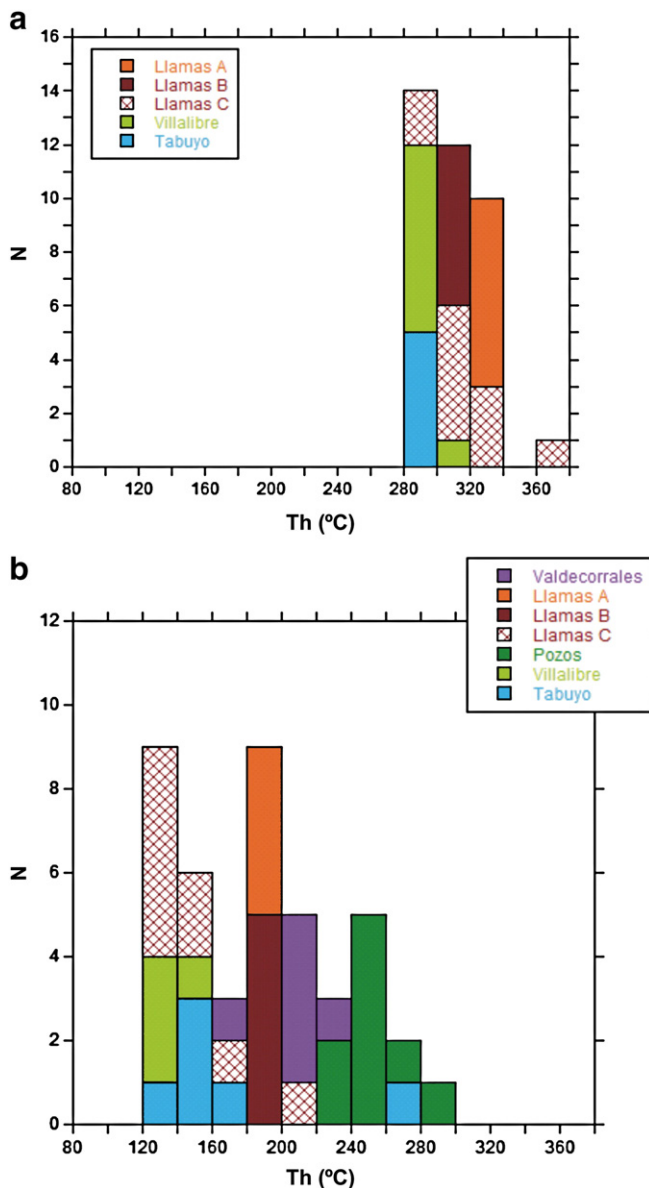


Fig. 6. Histograms of homogenization temperature: (a) aqueous-carbonic fluid inclusions (b) aqueous inclusions.

(Fig. 9) shows the general fluid evolution, indicating a clear increase in fluid salinity with a decrease in T_h . The H_2O -NaCl- CO_2 -(CH_4) fluid could be interpreted as an early fluid coexisting with the formation of quartz that displays a low salinity range (1.4–4.1 wt.% NaCl equiv.) in a relatively high T_h range (285–360 °C). The late H_2O -NaCl fluid displays low to moderate salinity (2.4 to 8.6 wt.% NaCl equiv.) and lower homogenization temperature (130–210 °C). Fluids at Pozos deposit exhibit a peculiar evolution with dominantly H_2O -NaCl composition, showing higher homogenization temperatures (230–285 °C) compared to other mineralizations in the region and moderate salinities (4.7–7.1 wt.% NaCl equiv.).

The CO_2 -(CH_4) fluid inclusions with no visible water have not been observed before in the gold-quartz veins from the northwest of the Iberian Peninsula. Previous studies of gold mineralizations from this area (Boiron et al., 1996; Couto et al., 1990; Noronha et al., 2000; Tornos et al., 1997) or central Iberia (Murphy and Roberts, 1997) recognized only aqueous-carbonic fluids. CO_2 -rich fluid inclusions are ubiquitous in vein-quartz gold deposits hosted in metamorphic terranes as a subordinate population in close association with co-genetic aqueous-carbonic fluids (Klein et al., 2006). Several models

Table 5
Sulfur isotope data for sulfide minerals from main deposits of Llamas de Cabrera district.

District	Vein location	Sample	Mineral	$\delta^{34}\text{S}$ (%V-CDT)
Pombriego		PMB-31	Arsenopyrite I	12.7
		PMB-31	Pyrite I	14.2
		PMB-31	Pyrite I	13.9
Valdecorrales		VDC-47	Arsenopyrite I	8.0
Llamas de Cabrera	Zone A	LLA-1	Arsenopyrite II	11.1
		LLA-1	Arsenopyrite II	12.3
	Zone B (El Zanjón)	LLB-5	Arsenopyrite I	9.7
		LLC-131	Arsenopyrite I	13.0
		LLC-3Q	Arsenopyrite I	9.7
	Zone C	LLC-3C	Arsenopyrite I	10.2
		LLC-33	Pyrite II	11.2
		LLC-33	Arsenopyrite II	9.0
		LLC-8	Arsenopyrite II	9.5
Villalibre		VL-1	Arsenopyrite II	23.1
Pozos		PZ-2	Pyrite II	19.5
		PZ-4	Arsenopyrite II	13.0
		PZ-5	Sphalerite	11.0
Tabuyo		TA-40	Arsenopyrite I	16.3
Post-kinematic pyrite "Luarca slate"		ALB-43S	Pyrite	26.3
		ALB-44S	Pyrite	7.4
		PMB-29S	Pyrite	20.4
		MAN-36S	Pyrite	16.2

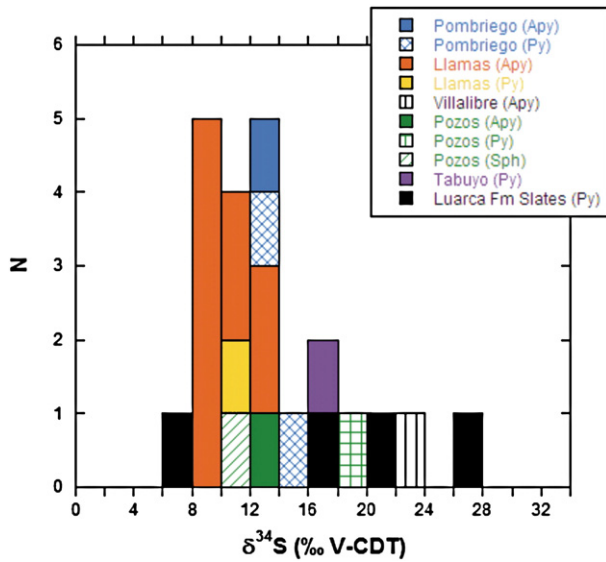


Fig. 7. Histogram of $\delta^{34}\text{S}$ for arsenopyrite, pyrite and sphalerite from main deposits of Llamas de Cabrera district.

have explained the origin of carbonic fluid inclusions in gold deposits (Chi et al., 2009; Garba and Akande, 1992; Wille and Klemd, 2004).

The absence of significant H_2O contents in these fluid inclusions contrasts with the formation of hydrous alteration minerals such as sericite and chlorite close to the veins. The CO_2 –(CH_4) fluid inclusions might be product of: (1) unmixing of a H_2O – CO_2 fluid before trapping of the inclusions; and/or (2) re-equilibration of early fluids during P–T-change or deformation during or after trapping. Unmixing could be a valid hypothesis if it occurs before trapping and outside the depositional site

Table 6

Oxygen isotope data for quartz. $\delta^{18}\text{O}_{\text{fluid}}$ correspond to the isotopic composition of water in equilibrium with quartz. Quartz calculations were performed for temperature from 285 to 330 °C and from using the fractionation equation of Zheng (1993).

District	Vein location	Sample	Mineral	$\delta^{18}\text{O}$ (‰ SMOW)	T_h (°C)	$\delta^{18}\text{O}_{\text{fluid}}$ (‰ SMOW)
Pombriego		PMB-31	Quartz II	13.7	325	7.4
		PMB-31	Quartz II	14.7	325	8.4
Valdecorrales		VDC-47	Quartz I	14.3	No data	
		VDC-47	Quartz I	14.4	No data	
Llamas de Cabrera	A Zone	LLA-1	Quartz I	9.3	330	3.2
		LLA-1	Quartz I	14.5	330	8.4
		LLA-1	Quartz II	14.3	330	8.2
	B Zone (El Zanjón)	LLB-1	Quartz I	14.6	300	7.6
		LLB-10	Quartz I	14.9	300	7.9
	C Zone	LLB-5	Quartz II	14.0	300	7.0
		LLB-5	Quartz II	13.7	300	6.7
		LLC-19	Quartz I	14.9	310	8.2
		LLC-2	Quartz I	15.1	310	8.4
		LLC-131	Quartz I	12.4	310	5.7
		LLC-131	Quartz II	16.1	310	9.4
		LLC-1011	Quartz I	14.1	310	7.4
		LLC-33	Quartz II	15.2	310	8.5
		LLC-11	Quartz II	14.0	310	7.3
LLC-11	Quartz II	14.5	310	7.8		
Villalibre	LLC-3Q	Quartz I	14.4	310	7.7	
	LLC-3Q	Quartz I	14.9	310	8.2	
	LLC-8	Quartz I	14.7	310	8.0	
	VL-1	Quartz I	14.9	285	7.4	
Pozos	PZ-2	Quartz I	15.1	No data		
	PZ-4	Quartz I	15.9	No data		
Tabuyo	TA-40	Quartz I	13.8	290	6.4	
	TA-40	Quartz I	13.8	290	6.4	

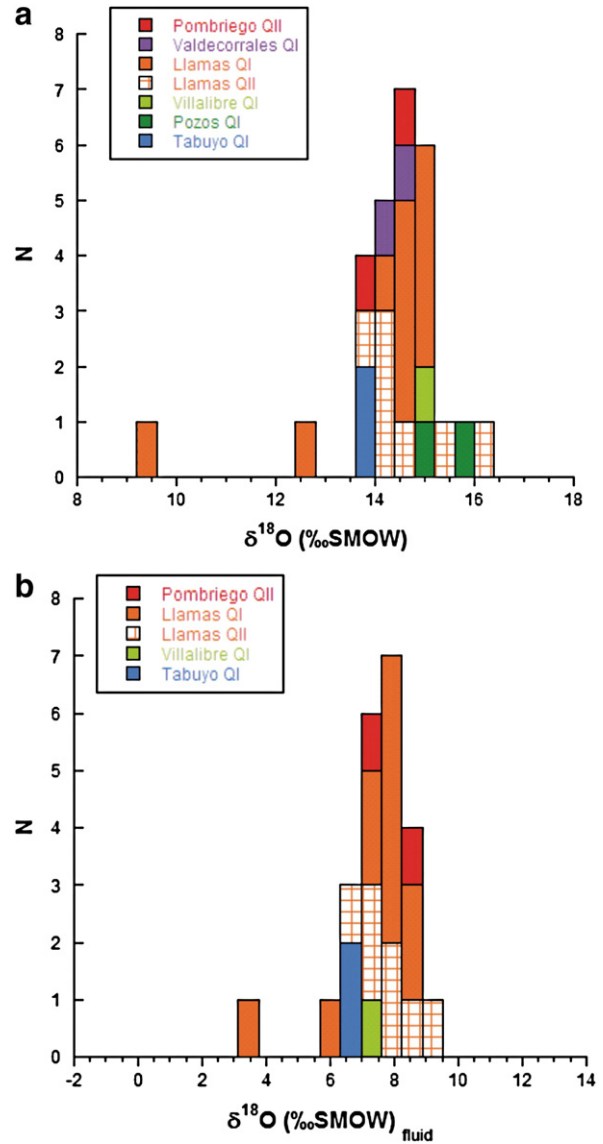


Fig. 8. a) Histogram of $\delta^{18}\text{O}$ values for quartz b) Calculated $\delta^{18}\text{O}$ fluid histogram.

(Chi et al., 2009, Klein and Fuzikawa, 2010). In this case the density and wetting contrast between CO_2 and H_2O favor trapping CO_2 -rich inclusions and the removal of the aqueous phase could contribute to the formation of hydrous alteration minerals in the vein-host rock contact. However, the presence of aqueous fluid inclusions along grain boundaries (Johnson and Hollister, 1995) and low trapping pressures would be expected. Trapping of fluid under low pressure is not consistent with the geological environment described for these deposits. Moreover, deformation could induce leakage of H_2O from aqueous-carbonic fluid inclusions in quartz (Bakker and Jansen, 1991; Hollister, 1990, Huizenga and Touret, 1999; Klein et al., 2006). Based on the presence of milky quartz crystals (QzI) with undulose extinction and serrate grain boundaries, the absence of water in the CO_2 –(CH_4) fluid inclusions could be interpreted by an effect of local re-equilibration provoked by deformation of the host quartz under fluctuating pressure conditions and deformation. However, CO_2 –(CH_4) fluid inclusions show relatively high densities in a close range (0.85–0.9 g/cm³) and H_2O leakage would lead to a variable density decrease in the remaining fluid. Unfortunately CO_2 –(CH_4) fluid inclusions are scarce and accordingly their origin remains uncertain.

Early aqueous-carbonic fluid inclusions were trapped in QzI and QzII quartz and are related to arsenopyrite and pyrite deposition. A drop in

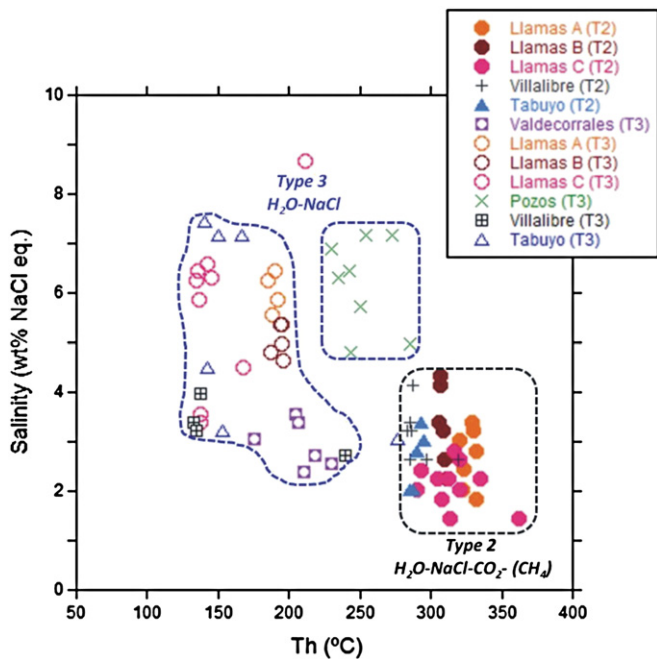


Fig. 9. Homogenization temperature versus salinity plot for $\text{H}_2\text{O-NaCl-CO}_2\text{-CH}_4$ (type 2) and $\text{H}_2\text{O-NaCl}$ fluids (type 3) in quartz.

temperature, decompression, and later stages of fracturing favored the infiltration of aqueous fluids which have been recognized along fluid inclusion planes. Aqueous fluids associated with gold paragenesis have low to moderate salinity (2.4 to 8.6 wt.% NaCl equiv.) and minimal trapping temperatures ranging from 130 to 285 °C. The T_h range is too large to be simply the result of a pressure change, indicating a mixing process between two end-members during cooling. The first one had a moderate salinity, up to 9 wt.% NaCl equiv., and the second one less than 2 wt.% NaCl. Fluid inclusions recognized in Pozos deposit could represent the relatively saline hot aqueous fluid (Fig. 9). A mixing process is consistent with the findings for comparable As–Au mineralizations elsewhere in the Variscan belt of western Europe (Boiron et al., 2003). In conjunction with arsenopyrite geothermometry and fluid inclusion data, an attempt was made to determine the ore-forming conditions. Because of its refractory nature, arsenopyrite does not re-equilibrate on rapid cooling and equilibrium compositions are usually preserved in hydrothermal deposits reflecting formation temperature (Choi and Youm, 2000). The arsenopyrite geothermometer for ore deposits formed at temperatures higher than 300 °C was proposed by Kretschmar and Scott (1976) and re-examined by Sharp et al. (1985). The average composition of arsenopyrite I in pyrite I assemblages clusters at 30.51–30.95 at.% As in Tabuyo deposit and at 28.37–30.07 at.% As in Llamas, corresponding to temperatures of 385 ± 10 °C and 330 ± 10 °C respectively. The average composition of arsenopyrite II in association with pyrite II exhibits a narrow variation, from 28.06 to 30.07 at.% As in Llamas, 28.47 to 30.03 at.% As in Pozos, and 28.76 to 29.93 at.% As in Pombriego. Temperatures estimated from these data are 320 ± 20 °C, 325 ± 20 °C, and 325 ± 15 °C (Table 7). Decreasing As contents of arsenopyrite indicate a decrease in sulfur fugacity or/and temperature during stage II. No differences in As values are detected between the different deposits during stage II. The arsenopyrite geothermometry provides temperatures slightly above homogenization temperatures of aqueous-carbonic fluid inclusions (285–330 °C with one outlier at 360 °C).

In a temperature range of 250–400 °C substantial amounts of gold may be transported in chemical conditions where bisulphide complex predominates. The decrease in gold solubility below 200–250 °C together with a pressure drop, cooling, pH change, H_2S loss (sulphidation

of iron-bearing minerals in the wall rocks), and oxidation appear to be effective mechanisms for gold precipitation (Gibert et al., 1998). In the context of Llamas de Cabrera district, dilution and mixing of two aqueous fluids of relatively contrasted salinity and temperature were favorable factors for decreasing gold solubility and could allow gold deposition. This model of gold deposition is similar to that observed in similar deposits in the Variscan belt (Boiron et al., 2003; Vallance et al., 2003).

A reconstruction of the pressure change during the ore stages has been carried out based on fluid inclusion isochores for each fluid circulation stage (Fig. 10). Extreme isochores give an estimate of the minimum pressure of trapping at a given temperature, and by considering other temperature and pressure indicators it was possible to place some constraints on the P–T conditions. Under brittle condition of deformation quartz veins are formed and aqueous-carbonic fluids observed in quartz contemporary with arsenopyrite and pyrite, probably records lithostatic pressures after the sealing of the fracture system. The minimal trapping temperatures (T_h) of these fluids vary between 285 and 360 °C and correspond to minimal trapping pressures (P_h) in the 190–240 MPa range. Constraints for the real trapping temperature and pressure are the following: (i) arsenopyrite geothermometry gives temperatures between 300 and 390 °C and (ii) peak of the regional metamorphism lower than 200–220 MPa (Díez Montes, 2007). As $\text{CO}_2\text{-(CH}_4\text{)}$ fluid inclusions could be interpreted by an effect of local re-equilibration their high density could not be representative of their trapping conditions (Vityk and Bodnar, 1995).

The latest fluids, which are associated with gold and Zn–Cu–Pb sulfides deposition are aqueous and were trapped at lower temperatures (minimum trapping temperatures in the range of 130–285 °C). Conditions of trapping are difficult to constrain due to the lack of mineralogical or other indicators at this stage, however they were probably trapped at pressures (≤ 200 MPa) similar or lower than those of the preceding stage (similar depth). These pressures indicate trapping temperatures ranging between 180 and 310 °C. A decrease in temperature, dilution and mixing of moderate to low salinity aqueous fluids could be related to gold deposition.

9.2. Source of ore-forming materials and fluids

The average value of $\delta^{34}\text{S}$ of minerals can represent that of the total sulfur in the hydrothermal fluid when the mineral association is simple (Ohmoto and Rye, 1979). The relatively tight clustering of the $\delta^{34}\text{S}$ (+9.0 to +14.2‰) in sulfides from quartz-gold veins can be interpreted to indicate that the fluid redox state was below the $\text{SO}_2/\text{H}_2\text{S}$ boundary and H_2S was the dominant reduced sulfur species in the fluids (Ohmoto and Rye, 1979). In H_2S predominant fluid, sulfide minerals precipitated from solution would exhibit $\delta^{34}\text{S}$ values similar to the $\delta^{34}\text{S}_{\Sigma\text{S}}$ value of the fluids. The similar sulfur isotope composition of stage I sulfides (from +8.0‰ to +16.3‰, Table 5) to that of stage II sulfides (+9.0‰ to +19.5‰, with one outlier at +23.1‰) suggests that both mineral stages share an identical or directly comparable sulfur source.

As shown in Fig. 7, the $\delta^{34}\text{S}$ values of sulfides separated from quartz-gold veins are similar to those of post- S_1 pyrites from siliciclastic Lluarca Formation exhibiting $\delta^{34}\text{S}$ values between +7.4 and +20.4‰ (with the highest value at +26.3‰). This feature may reflect that country rocks of quartz-gold veins have provided sulfur to the fluid system. The hydrothermal transport would tend to homogenize the $\delta^{34}\text{S}$, so that ores would display more constant values than the sources.

Similar $\delta^{34}\text{S}$ values (+6.7 to +18.0‰) were reported by Tornos et al. (1997) in gold deposits from Teleno area and by Arias et al. (1997) in the gold Navia Belt (+9.7 to +20.6‰), both hosted by Cambrian–Ordovician rocks in West Asturian Leonese Zone.

Quartz (Qz I and Qz II) crystallized from a CO_2 -bearing fluid with a $\delta^{18}\text{O}_{\text{fluid}}$ component clustering from +6.4 to 9.4‰ (Fig. 8). These

Table 7

Average and range (at.%) of As content and estimated temperature ranges (minimum–maximum values) from arsenopyrite geothermometry (Kretschmar and Scott, 1976) in arsenopyrite (Apy) from different mine sites.

Mine site	Sample	At.% As		Range	N ^b	Estimated temperature (°C)	Apy type
		Average	S.D. ^a				
Tabuyo	TA-40	30.73	0.31	30.51–30.95	2	385 ± 10	I
Llamas	LLC-132B	29.47	0.53	28.57–30.07	8	330 ± 10	I
Llamas	LLC-8	29.07	0.76	28.59–29.42	11	315 ± 15	II
Llamas	LLC-132	29.21	0.85	28.42–29.80	21	320 ± 20	II
Llamas	VDC-47(2)	29.19	0.81	28.06–29.96	9	320 ± 20	II
Pozos	PZ-5(3)	29.20	0.76	28.47–30.03	20	325 ± 20	II
Pombriego	PMB-31E	29.12	0.47	28.76–29.93	6	325 ± 15	II

^a Standard deviation.

^b Number of spot analyses by electron microprobe.

values lie within the range for a metamorphic source (Sheppard, 1986) and are characteristic of fluids isotopically in equilibrium with common basement materials at medium to high temperatures (285–330 °C). Taking the absence of igneous rocks in the Llamas de Cabrera area into account, calculated $\delta^{18}\text{O}_{\text{fluid}}$ values suggest that aqueous-carbonic fluids correspond to metamorphic (dehydration–decarbonation during prograde metamorphism) or aqueous fluids of unknown origin largely equilibrated with metamorphic lithologies between 300 and 400 °C.

The high temperature implies high heat flows which have been related in other gold deposits of the northwest of the Iberian Peninsula to the presence of granites (Boiron et al., 2003; Noronha et al., 2000; Vallance et al., 2003). In the case of Llamas de Cabrera the lack of igneous rocks exposed in the area requires some interpretation to explain the source of heat. Actually, there are several types of igneous rocks underlying the mineralized area (cross section, Fig. 1b) but most of them present characteristics which preclude a possible genetic relation to the gold ore. The volcanic related Ollo de Sapo Formation magmatism has been dated 472 to 488 Ma (U–Pb on zircon) by Díez Montes (2007); this means they are Upper Cambrian to Lower Ordovician in age and clearly pre-Variscan. Another group of intrusive bodies were classified as early syn-kinematic and syn-kinematic epi- and mesozonal granitoides; all of them underwent Variscan deformation and regional metamorphism with a peak and partial melting during D2. This metamorphism (M2/D2) was dated in the Ollo de Sapo antiform at 314 to 311 Ma by the U–Pb method on monazite. The peak P and T conditions (480–520 °C, 450–500 MPa) recorded in the Ordovician metasediments (Díez Montes, 2007) were much higher than those suggest here for the Llamas de Cabrera

Au mineralization. Thus, Au mineralization must have occurred later than M2. The only other magmatism known in the region is the Variscan post-kinematic Veiga granodiorite and other minor similar bodies placed to the east (Díez Montes, 2007). The Veiga granodiorite was dated by Ortega Cuesta et al. (2000a, 2000b) at 286 ± 6 Ma using the Rb/Sr method on whole rocks. These authors suggest that it is a post-kinematic epizonal body intruded during the extensional collapse which followed the Variscan crustal thickening. Furthermore, the estimated P and T grade at the Truchas Syncline during the late-Variscan extensional events is strongly consistent with the conditions presented here for the Llamas de Cabrera gold mineralization. Therefore, the ascent of some late Variscan (post-kinematic) magmatic bodies, not exposed at the present day erosion level, could act as a strong heat engine and remobilize gold from the metamorphic rocks. The age of this event (around 290 Ma) is younger than the metalliferous peak (310–300 Ma) established in the French Massif Central (Bouchot et al., 2005).

The deposition of gold is controlled by successive stages from high temperature aqueous-carbonic to low temperature aqueous fluid flowing along extension faults, breccias, and stockworks at the end of the Variscan orogeny. The aqueous-carbonic fluids result from a strong interaction with metamorphic rocks and characterized an early ore-stage with quartz, arsenopyrite, and minor pyrite deposition. An intense microfissuring of quartz and previous sulfides favored the circulation of aqueous fluids, and the mixing of aqueous fluids leads to gold and Zn–Cu–Pb sulfides precipitation. The ore fluids were mobilized during the Variscan doming generated for the syn-orogenic extensional processes and the related decompressive melting events.

10. Comparison with other gold Variscan deposits in western Europe

Gold-quartz veins from Llamas de Cabrera display similar features and processes to related deposits of similar type elsewhere in the Variscan belt of western Europe, northwest and central/western Iberia, the French Massif Central, and the Bohemian Massif (Table 8). The paragenesis reveals an early As–Fe stage and a consequent Pb–Zn–Cu–(Ag) stage, with Sb-rich mineral phases in some deposits. Native gold is associated with the second stage in most districts. Ubiquitous arsenopyrite generally has low gold concentration within its lattice. The presence of early aqueous-carbonic fluids that are associated with the As–Fe stage and that equilibrated with metamorphic rocks is a common feature for this mineralization type. The latest hydrothermal stage is characterized by the circulation of cooler aqueous fluids and by the mixing of moderately saline with dilute fluids of lower temperature. The $\delta^{18}\text{O}_{\text{fluid}}$ data obtained in northwest Portugal, Massif Central, and other gold districts in northwest Spain show similar values, characteristic of fluids isotopically in equilibrium with common basement materials, magmatic, or metamorphic rocks.

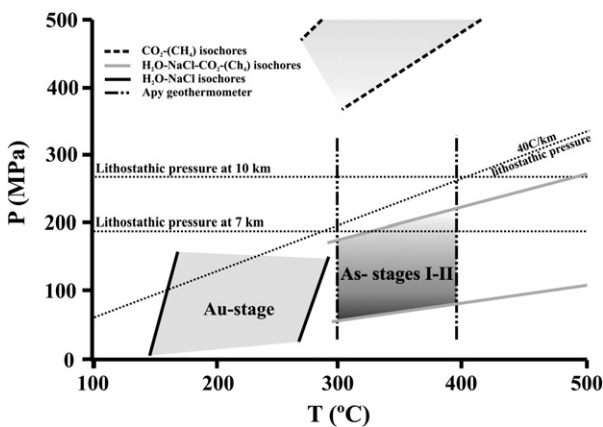


Fig. 10. Pressure–temperature reconstruction of conditions prevailing in Llamas de Cabrera deposit; data from mineral geothermometer and probable trapping conditions of the different of fluid inclusions.

Table 8
Characteristics of the main Variscan gold vein deposit in Europe.

Gold vein district	Orebody morphology and host rock	Paragenesis	Fluid composition and P-T conditions	Isotopic composition		Other characteristics	Age (Ma.)	References.
				$\delta^{18}\text{O}_{\text{fluid}}$ (‰ SMOW)	$\delta^{34}\text{S}$ (‰ V-CDT)			
Northwest Spain	Quartz–arsenopyrite gold veins in granites and Lower Cambrian to Middle Ordovician metasediments	Arsenopyrite–Pyrite–Gold \pm sulfides	CO ₂ –(CH ₄) H ₂ O–NaCl–CO ₂ –(CH ₄) P–T: 200–220 MPa, 300–390 °C H ₂ O–NaCl P–T: \leq 200 MPa, 180–310 °C	+ 3.2/+ 9.4	+ 6.7/+ 18.0	Stable isotope and fluid composition indicate fluids equilibrated with host rocks		This study Arias et al. (1997) Boiron et al. (1996) Tornos et al. (1997)
Northwest Portugal	Quartz–arsenopyrite gold veins in granites and Cambrian to Middle Ordovician metasedimentary rocks.	Arsenopyrite–Pyrite–Gold \pm sulfides	H ₂ O–NaCl–CO ₂ –CH ₄ P–T: 100–300 MPa, 300–500 °C H ₂ O–NaCl P–T: $<$ 100 MPa, $<$ 300 °C	+ 9.6/+ 11.4	No data	Gold precipitation linked to mixing to mixing between two distinctive aqueous fluids		Noronha et al. (2000) Vallance et al. (2003)
	Sb–Au veins in Cambrian metasedimentary rocks	Arsenopyrite–Pyrite–Stibnite–Gold \pm sulfides and sulfosalts	H ₂ O–NaCl–CO ₂ –(CH ₄) No P data, T: 300–390 °C H ₂ O–NaCl No P data, T: 125–225 °C	No data	No data	Pb isotope data indicate a homogeneous lead source of crustal origin		Neiva et al. (2008)
Central/Western Iberia	Syn and postcinematic Sb–Au veins in Upper Precambrian and Lower Palaeozoic metasedimentary rocks	Arsenopyrite–Pyrite–Stibnite–Gold \pm sulfides and sulfosalts	H ₂ O–(NaCl)–CO ₂ –CH ₄ –(N ₂) P–T: 150 MPa, 300 °C H ₂ O–NaCl No P–T data	No data	No data	Local unmixing associated to gold deposition		Dee and Roberts (1993) Murphy and Roberts (1997) Ortega and Vindel (1995)
Massif Central, France	Quartz–arsenopyrite gold–quartz lenses in Late Precambrian–Early Palaeozoic metamorphic rocks	Arsenopyrite–Pyrite–Gold (\pm sulfides)	H ₂ O–NaCl–CO ₂ –(CH ₄ –N ₂) P–T: 40–120 MPa, 260–550 °C H ₂ O–NaCl P–T: 40–120 MPa, 200–250 °C	0/+ 12.5	No data	Gold precipitation in relation to dilution and mixing of fluids, together a pressure drop	320–290	Boiron et al. (2003) Bouchot et al. (1989) Bouchot et al. (2005) Essaraj et al. (2001) Groves et al. (1998)
	Quartz veins hosted by Carboniferous migmatites and gneisses	Pyrite–Arsenopyrite–(Gold)–Stibnite \pm sulfides and sulfosalts	H ₂ O–NaCl (Trace of CO ₂ –CH ₄ – \pm N ₂) P–T: $<$ 5 MPa, 200–360 °C	– 1.0/+ 4	No data	Ore fluids characterized by an aqueous composition	310–300	Bellot et al. (2003) Boiron et al. (1989) Bouchot et al. (2005)
The Bohemian Massif	Quartz–arsenopyrite–gold veinlets in granodiorite	Arsenopyrite–Pyrite–Pyrrhotite–Gold (\pm minor sulfides)	H ₂ O–NaCl–CO ₂ –(CH ₄ –N ₂) P–T: 120–180 MPa, 350–450 °C H ₂ O–NaCl P–T: 20–70 MPa, 130–280 °C	No data	No data	Fluids of unknown origin but equilibrated with metamorphic rocks	340–310	Boiron et al. (2001) Groves et al. (1998) Morávek and Poucha (1990)

Although granites are present in the surroundings of most studied deposits, the granitoids act only as sources of heat that drive large convective hydrothermal cells along the major structures (Table 8).

Llamas de Cabrera deposits show certain differences in relation to other Variscan gold types, such as the presence of a CO₂–(CH₄) fluid in one deposit and the absence of an Sb stage; however the major difference is the absence of granites at the surface. The ascent of some late Variscan magmatic bodies, not exposed at the present day erosion level, could be invoked as the source of heat flow to sustain the hydrothermal activity. Another peculiarity of the Llamas de Cabrera veins is their consistent location in the transitional series called Rubianes Formation (quartzites, slates, and oolitic iron formation) overlying the Armorican quartzites. The reason for the relation between the gold-bearing quartz veins and this particular layer remains unknown.

As–Au hydrothermal veins such as Llamas de Cabrera are considered “deep-seated gold deposits” in the sense of Bouchot et al. (2005) developed under lithostatic to hydrostatic pressure. The majority of deposits of this style are sited in ductile to brittle structures and formed a depths from 6 to 12 km. Gold deposits of this type in the Variscan Iberian belt are restricted to the northwest of Spain and Portugal. In contrast, Sb–Au deposits are classified as “shallow gold deposits” emplaced from 6 km to near the surface under hydrostatic pressures. The shallow type includes several deposits in northwest/central Portugal and western Spain as syn- and post-kinematic veins (Table 8). The absence in Llamas de Cabrera of a Sb stage is probably due to deeper formation of the As–Au veins compared to other Sb–Au veins of northern, central, and western Iberia.

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