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ABSTRACTS











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and Early Permian. *C. tuberculata* was found in the Famennian of Poland and Artinskian of the Urals. *Laccophyllum* and *Caninia* were distributed everywhere, but *Laccophyllum turbinatum* and *Caninia kassini* were only found in the Tournaisian of Central Kazakhstan.

The rugosan complex from the "Imitoceras" layers of Berchogur is represented mostly by corals of the second morpho-ecological group (Kossovaya, 1994). However, Caninia of the first morpho-ecological group is also present.

The analysis of the rugosan complex from the "*Imitoceras*" layers shows the overpowering development of a Permian–Carboniferous coral association. *Laccophyllum* (Silurian–Early Carboniferous) is the only exception, which could be an effect of the mixed genus composition; so, the additional revision is necessary.

References

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THE FIRST REEFAL EPISODE (EARLIEST PRAGIAN) IN THE CANTABRIAN ZONE (NW SPAIN) AND ITS RELATIONSHIP WITH THE LOCHKOVIAN-PRAGIAN BOUNDARY EVENT

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In the Cantabrian Zone (NW Spain), as in other North Gondwanan regions, the first Devonian stage of reef development (just a few metres above the Lochkovian–Pragian boundary) is of minor importance and it is represented by small biostromal patches composed of branching rugose corals (disphyllids), branching tabulate corals (thamnoporids) and laminar to tabular, sometimes domical, stromatoporoids (coenostromatids, labechiids, atelodictyids, actinostromatids) and tabulate corals (favositids, micheliniids). These biostromal units are located in the middle part (Member C) of the Lebanza Formation (Palencia province; Palentine Domain, SE Cantabrian Zone) and in the upper part of the Nieva Formation (Asturias province; Asturo-Leonese Domain, N Cantabrian Zone).

The chronostratigraphic Lochkovian–Pragian boundary is located, as in Bohemia (where the event was defined with the name of the boundary), slightly below a marked colour change of the carbonate rocks, from dark to light, indirectly correlated with the *Eognathodus sulcatus* Zone (*Nowakia acuaria* Zone). Such a lithologic change was interpreted in different Devonian basins of the world as indicating a quick regression accompanied by a short-term anoxic event (Chlupác & Kukal, 1988).

The colour change coincides with a faunal turnover consisting of the decline of several brachiopods (Faunal Intervals 4–5 transition from García-Alcalde, 1996) and trilobite lineages. In addition, shortly afterwards, the first appearance of new benthic faunas (brachiopods, corals-stromatoporoids) occurs.

To sum up, the Lochkovian–Pragian Event is a geo-event of minor importance, but it is well represented worldwide (Saxonian-Thuringian and Frankenwald regions, Carnic Alps, Sardinia, NW Africa, Armorican Massif, Western and Eastern North America, Czech Republic and Spain), with substantial change in the faunal associations and appearance of the first reef episodes from the Lochkovian to Pragian.

References

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